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**IDI/SuperDARN
comparison**

R. E. Hibbins and
M. J. Jarvis

A long-term comparison of wind and tide measurements in the upper mesosphere recorded with an imaging Doppler interferometer and SuperDARN radar at Halley, Antarctica

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Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Abstract

Data from a co-located imaging Doppler interferometer and SuperDARN radar recorded since 1996 have been analysed in a consistent manner to determine daily mean winds and tides in the upper mesosphere. By comparing only days when both techniques were recording good quality data it is shown that the SuperDARN radar winds and tides correlate best with the IDI height bin 90–95 km. On timescales of one hour the winds derived from each technique correlate poorly, whereas the daily mean winds are in much better agreement suggesting that the two radars are sensitive to different parts of the gravity wave spectrum. Regression analysis reveals that the observed SuperDARN daily mean meridional wind strength is approximately 65% that recorded by the IDI while the zonal winds are of similar magnitude, in good quantitative agreement with previous studies which have shown contamination to SuperDARN-derived winds due to the significant back lobe of the radar radiation pattern. Climatologically the two techniques observe similar monthly mean winds with the SuperDARN meridional winds suppressed compared to the IDI which tends to record winds more poleward and eastward than those derived by the SuperDARN radar during the summer months, and to be slightly more equatorward during the winter. The 12-h tidal amplitude and phase in both the zonal and meridional components derived from both techniques are in excellent agreement, whereas the 24-h tides are seen much more strongly in the SuperDARN radar, especially in wintertime, with poor phase agreement. Long term comparison of the two techniques reveals a tendency for the IDI meridional winds to be more poleward during solar maximum especially during summer time; an effect which is not reproduced in the meridional winds derived from the SuperDARN radar. These results are discussed in the context of previous studies to independently determine the veracity of each technique.

IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

1 Introduction

The dynamics of the mesosphere and lower thermosphere (MLT) are dominated by gravity waves, tides and planetary waves. These periodic motions with timescales ranging from a few minutes to several months interact to drive the atmosphere to seasonal states far from radiative equilibrium especially in polar regions. Ground-based multi instrument studies are increasingly important in characterising the dynamics of the atmosphere through their inclusion in empirical models (e.g. Hedin et al., 1996; Portnyagin et al., 2004) and in deriving the spatial characteristics of tides and planetary waves (e.g. Murphy et al., 2006; Baumgaertner et al., 2006; Espy et al., 2005) to determine the geographical extent of different components. In Antarctica there are a mixture of different ground based radar techniques for studying winds in the mesosphere and lower thermosphere that can be used to separate and characterise wave components and the mean flow for climatological and phenomenological studies including imaging Doppler interferometer (e.g. Jones et al., 1997), MF spaced antenna (e.g. Jarvis et al., 1999), VHF all sky meteor drift (e.g. Forbes, Makarov and Portnyagin, 1995) and SuperDARN (e.g. Jenkins et al., 1998) radars. Cross calibrating different techniques is a vital part of this work and requires long-term high quality data sets from near or co-located systems.

At Halley (76° S, 27° W) an imaging Doppler interferometer (IDI) has been running since December 1996 recording winds between ~75 and 105 km, and a climatology of winds and tides has recently been published (Hibbins et al., 2006). At the same site a SuperDARN radar has been running since 1988 primarily recording coherent scattering from F-region plasma irregularities, but also has been used to derive winds in the lower thermosphere by tracing the ionisation trail left by meteors as they ablate in the MLT.

In this paper mean winds and tides are derived in a consistent manner from both techniques. Where the two techniques simultaneously generate data of consistent high quality comparisons are made between them. Section 2 covers the two data sets and

IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

the analysis techniques used to derive the mean winds and tides and their correlations on various time scales. The next section compares the mean winds and tides derived from the two techniques climatologically and quantifies the observed differences. Section 4 compares each wind component to look for systematic biases in one technique compared to the other and Sect. 5 presents evidence for long-term changes in the wind field recorded by each technique. Finally the differences are summarised in the light of previous attempts to calibrate these two techniques against other instruments.

2 Data

2.1 Instruments and data coverage

The IDI at Halley (Jones et al., 1997), based upon the technique of Adams et al. (1985, 1986) is a modification of the Halley Dynasonde, a NOAA HF radar (Grubb, 1979; Wright and Pitteway, 1979), and its operation has been described in detail elsewhere (Jones et al., 1997; Charles and Jones, 1999; Hibbins et al., 2006). The Halley SuperDARN radar (Greenwald et al., 1985, 1995) has been operating since 1988 primarily recording F region convection from coherent scatter off plasma irregularities in the ionosphere. Previous authors have demonstrated that the grainy near range echoes recorded by these radars are due to the ionisation trail left by meteors as they ablate in the upper mesosphere and lower thermosphere (Hall et al., 1997), and tracing the evolution of these trails provides an estimate of the neutral wind at these altitudes.

Figure 1 shows the data coverage recorded by the two radars between 1996 and 2006. Over this period the SuperDARN radar has generated hourly mean winds for 84% of the time with minor periods of down time for system maintenance primarily during the austral summer. For the IDI 67% of hours since December 1996 have generated wind data, but between 22 January 2002 and 18 February 2003 the IDI soundings were made on a delta antenna while the log-periodic antenna used for routine IDI soundings at 2.75 MHz was repaired following damage; this resulted in a 60% decrease in the

IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

number of returned echoes and consequently much poorer quality winds (Hibbins et al., 2006). These data have therefore been excluded from this study resulting in 54% of the time when both radars are simultaneously recording good quality data.

Figure 2 shows the normalised diurnal and annual variation in number of echoes recorded with the SuperDARN radar in range gates 0–4 (representing line of sight distances between 180 and 405 km recorded in all beams covering a $\sim 50^\circ$ horizontal field of view) and with the IDI in the altitude bin 90–95 km. Both techniques show similar strong diurnal and annual variations with minima in the number of recorded echoes around equinox especially in spring time. The SuperDARN radar, in common with other meteor radars, shows the greatest number of echoes in summer time around 07:00 LT ($\sim 09:00$ UT) whereas the IDI at this altitude records the greatest number of echoes in mid winter between 03:00 and 08:00 UT. The SuperDARN radar also tends to show a stronger diurnal variation in number of echoes than the IDI typically varying by around a factor of 3 between the maximum at 07:00 LT and the minimum around 19:00 LT.

2.2 Derivation of hourly mean winds

For the IDI hourly mean zonal and meridional winds were calculated from the three-component winds recorded every 5 or 15 min for each 5 km altitude bin between 75 and 105 km provided at least two different data points were present in the height bin during that hour. For the SuperDARN radar two-component horizontal hourly-mean winds were derived by singular value decomposition (Press et al., 1992) of all echoes with a signal to noise greater than 3 dB returned from the five range gates between 180 and 405 km from the radar assuming a scattering height of 95 km. Line of sight velocities greater than 100 ms^{-1} , or with an error greater than 50 ms^{-1} , or a spectral width greater than 25 ms^{-1} , were excluded from the fitting, and an hourly fit was only attempted if echoes were reported in at least five different radar beams. The zonal component of the resolved SuperDARN winds is seen to be noisier than the meridional component as expected from the observing geometry of the Halley radar. The boresite of the Halley

IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

radar is approximately 15° east of the geographic south pole and hence the Doppler velocities seen in the individual beams influence the resolved wind in this direction most strongly. Conversely the perpendicular winds are derived from differences in the radial velocities across the 16 radar beams spread out over only 52°.

Figure 3 shows the correlation coefficients between the hourly-mean meridional winds derived from both radars as a function of time of the day and month of the year. The correlation coefficients are seen to roughly follow the seasonal and diurnal variations in the number of echoes recorded by the two radars with a better correlation in winter than the equinoxes and a minimum between 12:00 and 21:00 UT. However, the magnitudes of the correlation coefficients are small with a maximum of ~0.3 between 06:00 and 09:00 UT in July, but virtually zero correlation in the afternoon and around the equinoxes. It is unlikely that the small geographical difference between the fields of view of the two radars will have a significant effect on the observed magnitude of the tides (see Sect. 3) therefore the poor correlation in the mean winds observed between the two techniques on timescales of 1 hour suggests the two techniques are observing different gravity wave fields. Whether this is due to the geographical difference between the two radars' field of view or to the sensitivity of the two radars to a different spectrum of waves is undetermined.

2.3 Tidal fitting

In order to separate the influence of the tides from the mean wind field, each horizontal component of the hourly-mean winds from each radar were combined into four-day segments and a mean bias and 12- and 24-h sine wave representing the two strongest tidal components (Hibbins et al., 2006) were fitted to each data segment provided at least half of the hourly-means were present spanning at least 16 different hours. Each individual hourly-mean was weighted in the fitting routine by the number of echoes recorded in that hour. In the case of the IDI each individual 5 km height bin between 75 and 105 km was treated independently in the fitting. The four-day segments of data were stepped by one day intervals to build up a time series of background winds and

IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

12- and 24-h tidal amplitudes and phases (as measured by the local time of maximum northwards or eastwards wind) recorded by the two radars since 1996.

2.4 Correlation with height

Figure 4 shows the reduced Pearson correlation coefficient for the meridional mean wind and the 12- and 24-h tidal amplitudes and phases between the SuperDARN radar and each of the 5 km height bins of the IDI. It can be seen that the mean wind recorded with the SuperDARN radar correlates best with the IDI mean wind recorded in the 90–95 km height bin consistent with the mean height of 94(±3) km quoted by Hall et al. (1997). Though each 5 km IDI height bin is oversampled by a factor of three (Jones et al., 1997) so the 90 km bin actually cover a region between 85 and 100 km with a Gaussian envelope. The 12- and 24-h tidal amplitudes also agree with this correlation with coefficients peaking at the 90–95 km altitude bin. The 24-h tidal amplitudes correlate much more poorly than either the mean meridional winds or the 12-hour tidal amplitudes across all altitude ranges observed by the IDI, and the phases show no significant change in correlation with altitude which is expected as the diurnal tide at Halley has been shown previously to be largely evanescent (Hibbins et al., 2006). There is no seasonal variation in the height of the maximum of the correlation coefficients when the data are separated into calendar months, suggesting that the height of the meteor layer does not vary significantly (i.e. by more than 5 km) between polar winter and summer. For the rest of this paper it is assumed that the SuperDARN radar winds are observed from a region that agrees best with the IDI 90–95 km height bin.

2.5 Correlations with season

Figure 5 shows the correlation coefficients between the meridional mean winds and tidal amplitudes observed in the SuperDARN and IDI (90–95 km height bin) separated into calendar months. The tidal-subtracted daily-mean winds correlate better than the hourly-mean winds but with a similar peak in the degree of correlation around late win-

IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

ter coincident with the wintertime maximum in the number of echoes recorded by both radars. The semidiurnal tide shows the opposite behaviour peaking in summer time. This coincides with the period of largest tidal amplitudes (Hibbins et al., 2006). The fitted diurnal tidal amplitudes measured in the two instruments correlate much more poorly than the semidiurnal tides with virtually zero correlation around the equinoxes when the tides are weak and the overall wind quality is poorest. It would appear that on timescales of a few days the correlation in tidal amplitudes maximises when both the amplitudes are large and the background winds are derived from the largest number of echoes.

3 Climatologies

To further investigate the source of the differences in the fitted background mean winds and tides, climatologies of the monthly mean wind and 12- and 24-h tidal amplitude and phases for the meridional and zonal components of the SuperDARN meteor winds and the IDI winds measured in the 90–95 km height bin were generated and are presented in Figs. 6a and b. To aid comparison, only days where winds and tides were successfully fitted to the data generated by both radars are used in generating the monthly means, and the tidal amplitude and phases are vector averaged. Typically around 200 days of data are included in generating each data point. Included in the figures are the mean winds and the amplitudes and phases of the 12- and 24-h tides derived from the HWM-93 horizontal wind model (Hedin et al., 1996) assuming an altitude of 92.5 km and geographic locations of 76° S, 27° W (the IDI location) and 78° S, 24° W (the mean position of the scattering points measured in the first five range gates of the SuperDARN radar) to demonstrate the magnitude of the anticipated differences in the climatologies due to the different field of view of the two radars.

IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

3.1 Mean winds

Between March and September the meridional mean winds measured in the 90-95 km height bin of the IDI are more equatorward than those recorded by the SuperDARN radar, and more poleward during the summer months. The difference is more pronounced in December where the SuperDARN radar mean wind is nearly 6 ms^{-1} more equatorward, but differ by no more than 3 ms^{-1} during the equinoxes and winter. Similarly, the zonal winds differ most in December where the SuperDARN radar reports winds 8 ms^{-1} more westwards than the IDI, but between February and October the two techniques track each other closely differing by no more than 4 ms^{-1} . Both radars generate less seasonal variability in the zonal winds than predicted by the HWM-93 empirical model which shows eastwards winds of up to 10 ms^{-1} in winter and -5 ms^{-1} in summertime.

3.2 Tides

The diurnal tide derived from the two different techniques is significantly different. The amplitude of the meridional component is seen more strongly in the IDI in the summer months and more strongly in the SuperDARN radar at other times, a similar pattern is observed in the zonal component except the wintertime differences are much larger with the SuperDARN radar seeing tidal amplitudes up to 10 ms^{-1} larger than the IDI. The phases of the derived mean diurnal tides also differ especially during winter in the meridional component and summer in the zonal component. Despite the differences in the observed diurnal tides recorded by the two techniques, surprisingly the semidiurnal tides are almost identical. The amplitudes of both components track each other to within the standard errors throughout nearly all of the year, both showing stronger tides than those predicted by the HWM-93 model during summer when the non-migrating $S=1$ component of the semidiurnal tide becomes dominant at Halley latitudes (Baumgaertner et al., 2006). The phases are also in good agreement between the two techniques, only differing by a significant amount around the equinoxes in the zonal com-

IDI/SuperDARN
comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

ponent. The poor agreement in the diurnal tides compared to the semidiurnal tides is probably related to the diurnal variation in the number of echoes (and hence the quality of the hourly mean winds) seen throughout the year. The SuperDARN radar typically sees many more meteor echoes between 00:00 and 12:00 UT at Halley with diurnal variations of up to a factor of four in summer time. The IDI also sees a diurnal cycle in wind quality with around twice the number of echoes recorded between 22:00 and 10:00 UT than 10:00 to 22:00 UT in winter and the equinoxes, and a peak slightly later in the day in mid summer. Thus in any 24-h period the hourly mean winds recorded during one 12-h period will be much noisier than another. This should have little effect on 12-h tides fitted over a four day period but could bias the fitting of 24-h tides.

4 RMA regression

To determine if there is a tendency for one technique to measure systematically stronger winds than the other, a reduced major axis (RMA) regression analysis (Sokal and Rohlf, 1981) was undertaken on the daily-mean zonal and meridional winds, linear regression being inappropriate in this case as it assumes one independent variable with zero associated error and hence the gradient of any fit will tend towards the axis on which the independent variable is plotted. To illustrate this Fig. 7 shows a scatter plot of the daily zonal and meridional mean winds recorded by the two techniques. Overplotted are the least squares fit lines assuming linear regression with each technique as the assumed independent variable, and the RMA regression assuming the errors in each technique are similar. The fits show that the SuperDARN meridional winds are systematically weaker than those recorded by the IDI but the zonal winds are of similar magnitude with fits to the data giving straight lines of the form:

$$V_{SD} = 0.62(\pm 0.02) \cdot V_{IDI} + 0.13(\pm 0.13)$$

$$U_{SD} = 1.12(\pm 0.03) \cdot U_{IDI} - 2.45(\pm 0.25)$$

where V and U are the meridional and zonal winds respectively and the subscripts SD and IDI refer to the SuperDARN and IDI (90–95 km height bin) radars. The weak-

IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

ness of the SuperDARN meridional winds is likely to be due to the significant back lobe of the SuperDARN radar which is a mirror image of the forward radiation pattern (Milan et al., 1997) and at Halley points approximately equatorwards. Arnold et al. (2003) have shown that the meridional winds recorded in the Finland CUTLASS radar can be reduced by between 20% and 50% due to contamination of the wind field by meteors detected in the radar's back lobe depending on the operating frequency, time of day and season with the largest effect seen at the lowest frequencies when the meteor echoes are strongest. These values are in good agreement with the 38% reduction in meridional winds seen in the Halley SuperDARN radar compared with the IDI over the ten years of operation. This observation is further supported here when the RMA regression analysis is repeated separating the data into summer and non-summer months. During the summer months of November to February inclusive, when the SuperDARN echoes are strongest the RMA gradient is $0.55(\pm 0.03)$ whereas outside these months when the echoes are typically weaker the gradient increases to $0.69(\pm 0.02)$. Hence the SuperDARN meridional winds are closer in magnitude to the IDI winds when the echoes are weaker and the contribution from meteors detected from the radar's back lobe is lessened.

5 Long term change

The 11 years of data presented here span the complete cycle 23 of solar activity presenting an opportunity to search for deviations from the climatological mean in the winds and tides caused by variations in solar flux. Figure 8 shows a plot of the monthly mean deviation from the climatological mean in the meridional wind in both the IDI and SuperDARN radars versus the solar activity proxy measured by the monthly mean observed flux of solar radiation at a wavelength of 10.7 cm. Although the scatter in the data are greater, the IDI meridional winds clearly show a more polewards tendency at periods of high solar activity that is not reproduced in the SuperDARN meridional winds. The gradient of a linear least squares fit gives a mean variation

IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

of $-2.21(\pm 0.90) \times 10^{-2} \text{ ms}^{-1}/\text{SFU}$ corresponding to a $3.7(\pm 1.5) \text{ ms}^{-1}$ difference over a typical 170 SFU solar cycle, whereas this trend is not reproduced in the SuperDARN monthly meridional wind field over the same time period where a similar linear regression yields an insignificant gradient of $-1.10(\pm 42.0) \times 10^{-4} \text{ ms}^{-1}/\text{SFU}$.

Detailed discussions of these long term trends are beyond the scope of this paper and will be presented elsewhere, though Table 1 summarises the magnitude of the gradient of a linear fit to a plot of the perturbation in the monthly-mean winds and tidal amplitudes and phases against the monthly-mean index of solar activity. Both techniques clearly see a reduction in the zonal tidal amplitudes with increasing solar activity which is particularly present in the semidiurnal tide; this is around $2.8(\pm 1.5) \text{ ms}^{-1}$ weaker at solar maximum than at solar minimum with no significant change in phase. The strong poleward tendency for the IDI mean winds at solar maximum is not reproduced in the SuperDARN meteor winds. As comparisons are averaged over a whole cycle of solar activity it is clear comparison of the two techniques during periods of high solar activity will draw different conclusions from comparisons undertaken during solar minimum. To emphasise this point Fig. 9 shows a meridional wind climatology as presented in Fig. 6 for the IDI and SuperDARN radars derived from two separate time periods representing typically high solar activity (1999–2002) and typically low solar activity (1996–1998 and 2003–2006). The two climatologies generated from the SuperDARN meteor winds are both similar, apart from a single anomalous point in September, showing the insensitivity of the meridional winds to solar cycle. However the IDI climatologies show that at high solar activity the meridional winds are significantly more poleward throughout the summer months of October through to February. Comparison with Fig. 6a shows that the monthly mean meridional winds derived from the two techniques agree much better throughout the year during low solar activity.

**IDI/SuperDARN
comparison**

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

6 Discussion

A number of authors have previously compared winds derived from both IDI and SuperDARN radars with other methods of determining horizontal winds in the MLT. Typically the IDI is seen to generate somewhat weaker winds than comparable techniques. Turek at al. (1998) compared IDI winds between 70 and 100 km altitude with a co-located incoherent scatter radar from Arecibo for a 2 h period on 10 April 1989. Whereas previously Hines et al. (1993) had questioned the reliability of the IDI winds in this data set above 80 km altitude, they found good agreement between the two techniques except in the region of maximum wind shear above 90 km where the meridional winds were seen to be much weaker in the IDI derived winds, an effect ascribed to interference from gravity waves preferentially propagating in the north-south direction. On timescales of one hour we find relatively poor correlations between the winds recorded by the IDI and the SuperDARN radar probably due to the sensitivity of the two techniques to different spectra of gravity waves. Jones et al. (2003), in a more extensive study, compared an IDI system with a VHF SkiYMET all-sky meteor radar equipped with height finding over a four month period between November 2000 and March 2001 at Bear Lake observatory (42° N, 111° W). They found qualitatively good agreement between the two techniques, though histograms of the difference between the hourly mean winds recorded by the two techniques generated distributions with an offset of up to 8 ms⁻¹ and standard deviations of up to 26 ms⁻¹ at 98 km altitude. However, between the overlapping height ranges of 80–95 km during this period when the semidiurnal tide was strong, they concluded that there was a tendency for the meteor winds to be up to 10% stronger than those derived by the IDI. This is in good agreement with the RMA regression comparison presented here between the zonal mean winds derived from the two techniques which are seen to be around 10% stronger in the SuperDARN radar than the IDI at 90–95 km.

Hall et al. (1997) compared SuperDARN meteor wind velocities from a single beam of the Saskatoon radar with those determined from a co-located MF radar. During both

IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

August and January 1994 both techniques correlated best when comparing the MF radar winds recorded around 90(± 10) km altitude, although the correlation coefficients were higher in summer than in winter. In addition they observed that the MF horizontal wind velocities at 97 km altitude were about 16% greater than those determined from the SuperDARN single beam during August 1994, a difference largely ascribed to the 200–300 km differing locations of the scattering points from the two radars. By comparing the phase of the semidiurnal tide in each technique during those months of 1994 when the phase was seen to change rapidly with height they observed a best fit in phase of the SuperDARN winds with the MF radar winds recorded at 94(± 3) km. Hussey et al. (2000) extended this work to observe winds with a MF and SuperDARN radar at Saskatoon and concluded “the two systems complement each other effectively”. Over a two week period in September 1997 they observed strong semidiurnal tides in both techniques and saw that the SuperDARN radar matched best with the MF radar at an altitude of 94–97 km based on the phase of the semidiurnal tide. In addition they concluded that the MF radar winds were approximately 10% smaller than those observed by the SuperDARN radar based on the means of the 12-h oscillations. This study also concludes that the semidiurnal tide is in good agreement between the two techniques with the SuperDARN radar generating a best fit to the IDI semidiurnal tides at an altitude of 90–95 km, but with the IDI typically producing slightly larger semidiurnal tidal amplitudes than those derived from the SuperDARN winds. Arnold et al. (2003), in a more extensive comparative study, derived meridional wind climatologies based on one year of data recorded between September 1999 and September 2000 from the Hankasalmi SuperDARN radar in Finland and a VHF meteor radar located around 700 km away in Kiruna, Sweden. They found best agreement between the two climatologies occurred from data recorded between 06:00 and 18:00 UT when the SuperDARN radar was operating at 12 MHz and contamination from the back lobe of the SuperDARN radar was minimised. Otherwise the SuperDARN winds were found to be up to 50% weaker than the VHF meteor radar winds, although a statistical variance test on the derived climatologies showed better agreement between the

**IDI/SuperDARN
comparison**R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

SuperDARN and VHF radars than was found by separating the VHF meteor echoes into north and south centred quadrants, suggesting a significant geographical variation in the observed winds. The data presented here also show that the SuperDARN meridional winds are suppressed with respect to the IDI winds by between 30 and 45% depending on the season, though these values are also shown to be dependent on the phase of the solar cycle due to the tendency for the IDI meridional winds to change with solar activity.

7 Summary

This paper has presented the first direct comparison of horizontal winds derived from a co-located imaging Doppler interferometer and SuperDARN radar, and the first comparison of any two MLT wind radars to span a complete solar cycle. We have demonstrated that correlations between the parameters derived from these two radars vary on timescales ranging from hours to years, and that a true assessment of the differences between the two techniques requires simultaneous co-located long-term data sets. Care must be taken when combining data recorded by different techniques to derive empirical models or to study the zonal structure and geographic extent of tides and planetary waves.

The relatively poor correlations between the hourly-mean winds is probably be due to differences in the observed gravity wave field. However, when the daily mean winds are separated from the influence of the tides, the winds derived from two techniques are seen to correlate much better, with the SuperDARN radar generating systematically weaker meridional winds due to the geometry of the radar's radiation pattern. Monthly mean climatologies generated by the two techniques show good agreement in the winds and semidiurnal tides, but very poor agreement in the derived diurnal tidal amplitudes and phases related to the strong diurnal influence in the returned echoes and hence the quality of the winds. Finally, interannual variations between the two techniques have shown that the meridional wind derived by the IDI is influenced by

IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

solar activity to a much greater degree than the SuperDARN winds, so much so that a climatology derived from IDI winds around solar maximum is significantly different from that recorded around solar minimum especially in summer time when the meridional winds recorded at periods of high solar activity are around 5 ms^{-1} more poleward than
 5 at low solar activity.

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IDI/SuperDARN comparison

R. E. Hibbins and
 M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

IDI/SuperDARN comparison

R. E. Hibbins and
M. J. Jarvis

Table 1. Gradient of the linear least squares fit between the monthly mean zonal and meridional winds and tides recorded by the IDI and SuperDARN radar and the monthly mean observed F10.7 cm solar flux. The numbers in parentheses represent 1σ errors on the fitted gradient. Cells where the calculated gradient is less than the error are shaded.

	zonal		meridional	
	IDI _{90-95km}	SuperDARN	IDI _{90-95km}	SuperDARN
mean wind (ms ⁻¹ /SFU)	-1.3(9.6)x10 ⁻³	-2.2(7.3)x10 ⁻³	-2.2(0.9)x10 ⁻²	-1.1(42)x10 ⁻⁴
24-h amplitude (ms ⁻¹ /SFU)	-4.2(1.2)x10 ⁻³	-8.3(7.3)x10 ⁻³	+4.2(7.9)x10 ⁻³	-2.9(4.2)x10 ⁻³
12-h amplitude (ms ⁻¹ /SFU)	-1.6(1.1)x10 ⁻²	-1.7(0.7)x10 ⁻²	-9.1(9.9)x10 ⁻³	-3.7(6.6)x10 ⁻³
24-h phase (hr/SFU)	-9.7(7.7)x10 ⁻³	-8.4(4.5)x10 ⁻³	-5.9(7.3)x10 ⁻³	+5.4(2.8)x10 ⁻³
12-h phase (hr/SFU)	+2.8(3.2)x10 ⁻³	-3.1(3.2)x10 ⁻³	+3.2(2.4)x10 ⁻³	-1.4(17)x10 ⁻⁴

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[◀](#)
[▶](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)

**IDI/SuperDARN
comparison**

R. E. Hibbins and
M. J. Jarvis

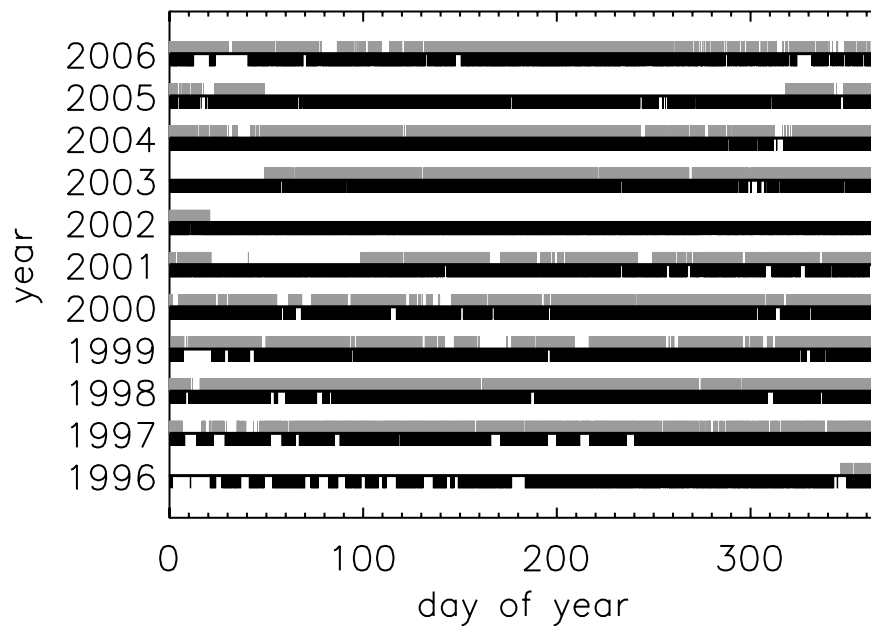


Fig. 1. Data coverage recorded by the SuperDARN radar (black line) and the IDI (grey line) between 1996 and 2006.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

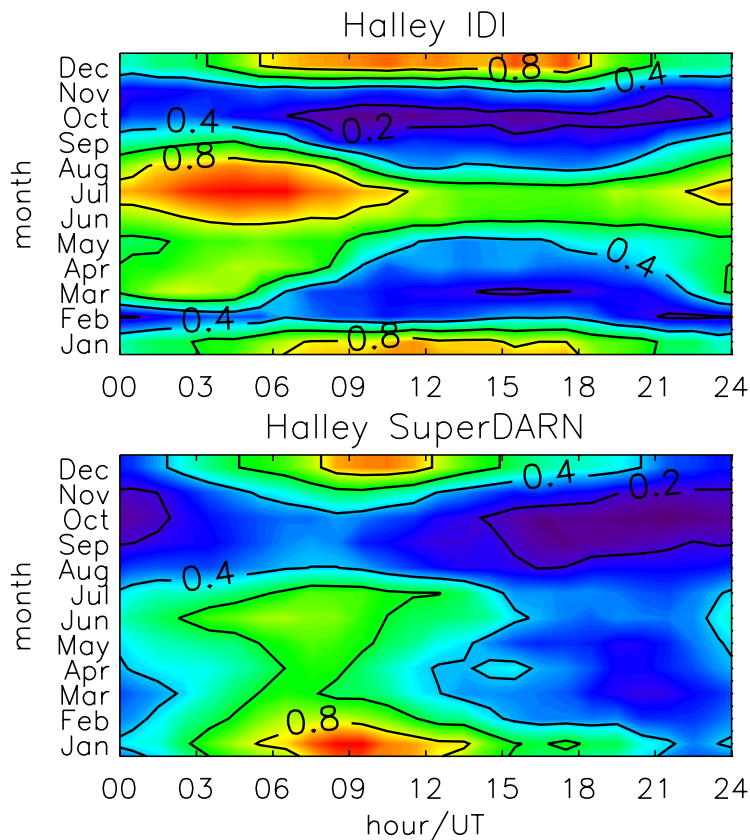
**IDI/SuperDARN
comparison**R. E. Hibbins and
M. J. Jarvis

Fig. 2. Annual and diurnal variation in the number of echoes recorded by the IDI 90–95 km height bin (top) and the SuperDARN radar (bottom). Both plots are independently normalised to emphasise the seasonal and diurnal variation.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**IDI/SuperDARN
comparison**

R. E. Hibbins and
M. J. Jarvis

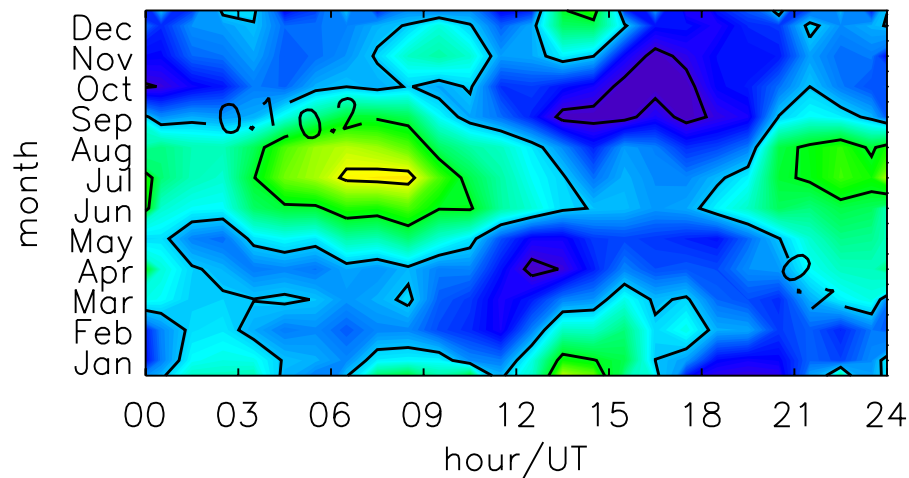


Fig. 3. Correlation coefficients between the hourly mean meridional winds recorded by the SuperDARN radar and the IDI 90–95 km height bin.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[I◀](#)[▶I](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**IDI/SuperDARN
comparison**

R. E. Hibbins and
M. J. Jarvis

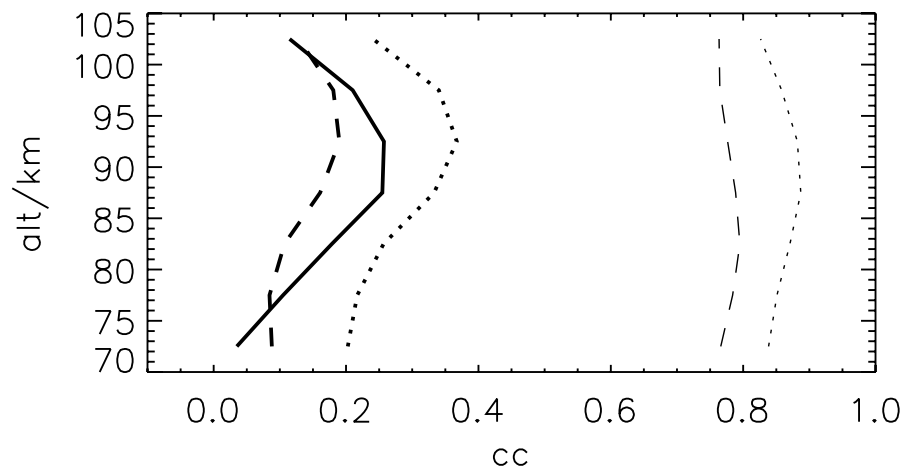


Fig. 4. Correlation coefficients between the daily mean meridional winds and tides recorded by the SuperDARN radar and the IDI for each 5 km altitude bin between 70 and 105 km. Thick solid line: mean wind, thick dashed line: diurnal tidal amplitude, thick dotted line: semidiurnal tidal amplitude, thin dashed line: diurnal tidal phase, thin dotted line: semidiurnal tidal phase.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**IDI/SuperDARN
comparison**

R. E. Hibbins and
M. J. Jarvis

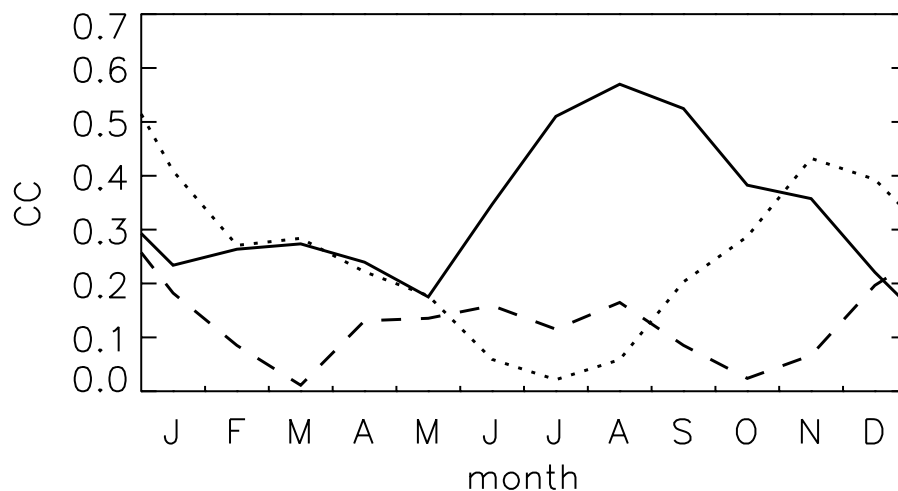


Fig. 5. Correlation coefficients between the daily mean meridional winds and tides recorded by the SuperDARN radar and the IDI 90–95 km height bin separated into calendar months. Solid line: mean wind, dashed line: diurnal tidal amplitude, dotted line: semidiurnal tidal amplitude.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

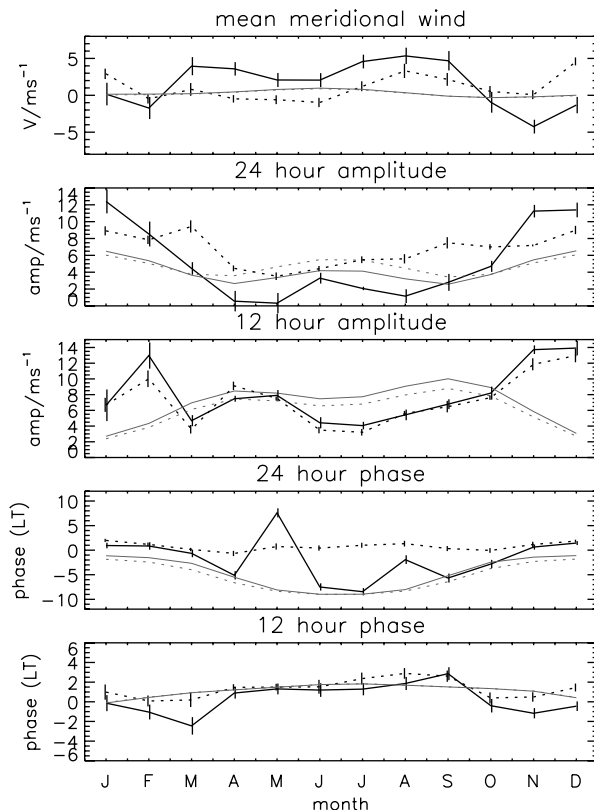
**IDI/SuperDARN
comparison**R. E. Hibbins and
M. J. Jarvis

Fig. 6a. Meridional wind and tide climatologies generated from the SuperDARN radar (dashed line) and IDI 90–95 km height bin (solid line). The thin grey lines represent the equivalent climatologies generated from the HWM-93 model (note: on some panels these overlap).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**IDI/SuperDARN
comparison**

R. E. Hibbins and
M. J. Jarvis

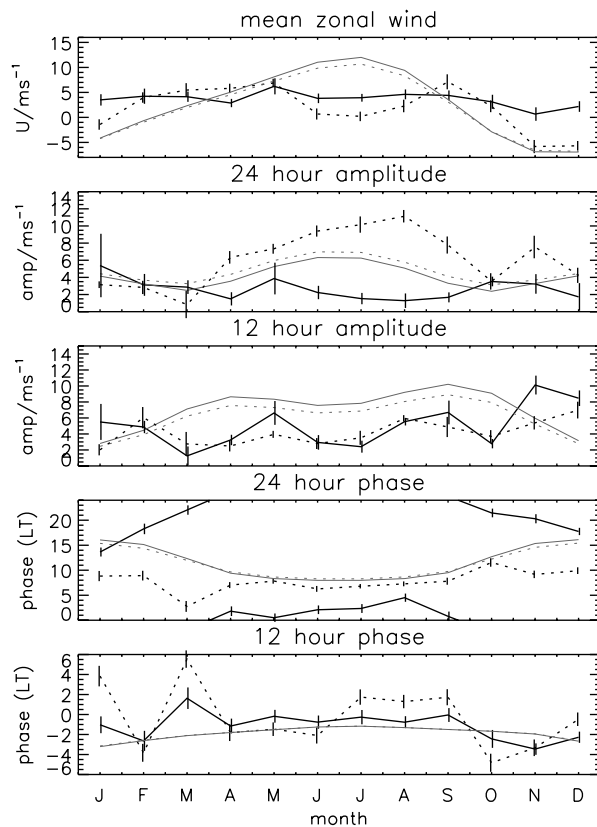


Fig. 6b. As for Fig. 6a but for the zonal component of the winds.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**IDI/SuperDARN
comparison**

R. E. Hibbins and
M. J. Jarvis

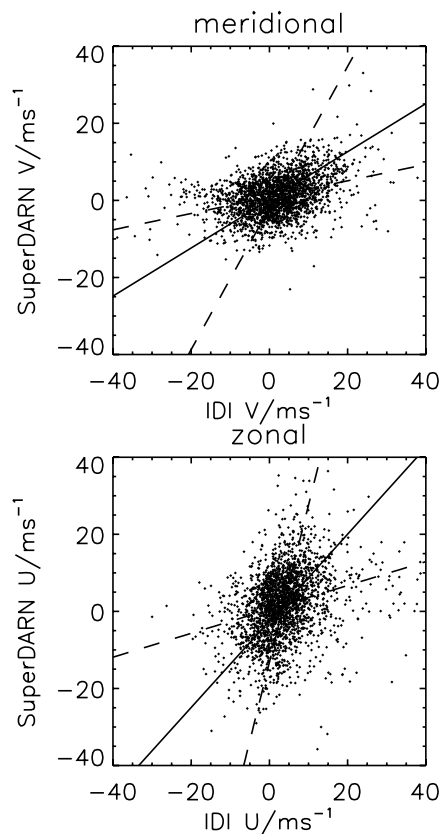


Fig. 7. Scatter plot of the daily mean winds recorded by the SuperDARN radar and the IDI 90–95 km height bin for the meridional (top) and zonal (bottom) components. Included on the plots are the RMS regression lines (solid) and the linear regression lines assuming each of the two variables is independent (dashed).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**IDI/SuperDARN
comparison**

R. E. Hibbins and
M. J. Jarvis

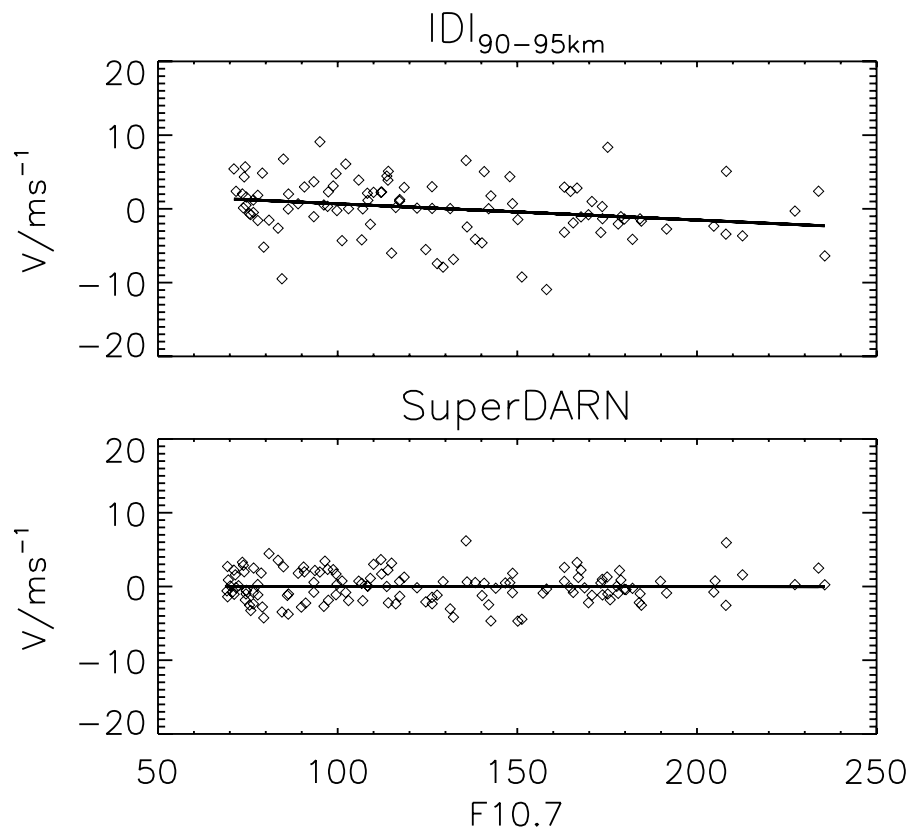


Fig. 8. Monthly mean meridional winds versus monthly mean F10.7 index for the IDI 90–95 km height bin (top) and the SuperDARN radar (bottom). Included on the plots are the lines of least squares fit.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**IDI/SuperDARN
comparison**

R. E. Hibbins and
M. J. Jarvis

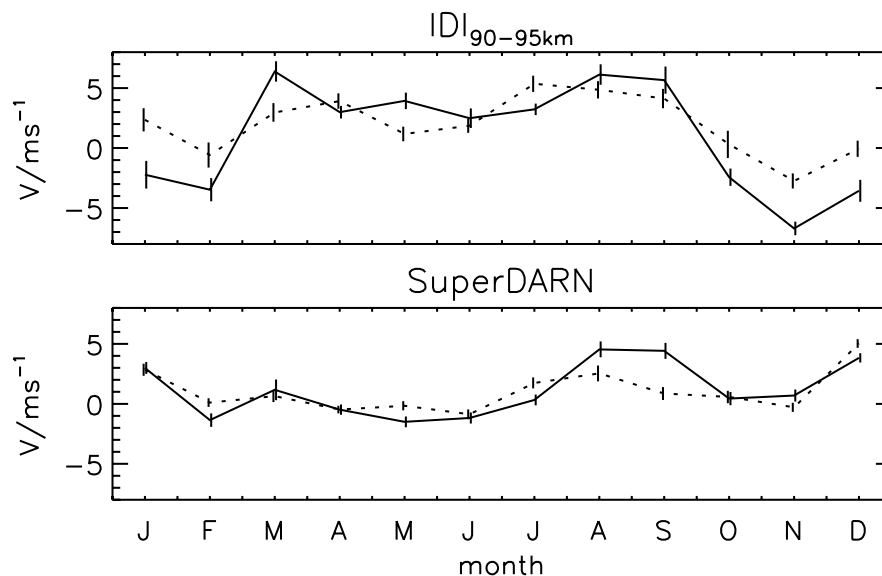


Fig. 9. Meridional wind climatology generated from the IDI 90–95 km height bin (top) and the SuperDARN radar (bottom). Solid line: data from 1999–2002 inclusive (high solar activity), dashed line: data from 1996–1998 and 2003–2006 inclusive (low solar activity).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)